#### Surface nutrient distribution



#### The 3 box model



Model equation: a system of 3 ODEs

$$V_L \frac{dP_L}{dt} = C(P_D - P_L) - B_L$$
$$V_H \frac{dP_H}{dt} = C(P_L - P_H) + M(P_D - P_H) - B_H$$
$$V_D \frac{dP_D}{dt} = (C + M)(P_H - P_D) + B_L + B_H$$

Model equation in vector-matrix formulation

$$\frac{d}{dt}\mathbf{P} = \mathbf{T}\mathbf{P}$$

# Solution methods

• Analytic

 $\mathbf{P}(t) = e^{\mathbf{T}t}\mathbf{P}(0)$ 

MATLAB

>> P(n) = expm(T\*t(n))\*P(1);

• Numerical (Euler forward)

$$\mathbf{P}_{n+1} = (\mathbf{I} + \Delta t\mathbf{T})\mathbf{P}_n$$

• Numerical (Euler backward)

$$\mathbf{P}_{n+1} = (\mathbf{I} - \Delta t \mathbf{T})^{-1} \mathbf{P}_n ~~$$
>> P(n+1) = (eye(3) - dt\*T)<sup>-1</sup>\*P(n);

# Practical recommendations

- Euler forward scheme can be unstable. Choose Euler backward if the situation allows (it requires matrix inverse calculation).
- If analytic solution is available, use it to check your numerical solution.
- Despite my earlier suggestion, Euler forward scheme can often produce acceptable results if a short enough timestep is used.

```
week10_exercise.m
 % PO4 box model
 % clean first
 clear all;
 close all;
 % 1. set up model parameter
 H = 4000; % full ocean depth, m
 R = 6.3e6; % radius of the Earth, m
 Area = 4*pi*R^2*0.8; % total surface area, m2
 fL = 0.8; % fraction of low lat ocean
 V(1) = Area*fL*100; % surface low lat ocean vol, m3
 V(2) = Area*(1-fL)*300; % surface high lat ocean vol, m3
 V(3) = Area*H-V(1)-V(2); % deep ocean vol, m3;
 C = 10*1e6; % circulation rate (m3/s);
 M = 40*1e6; \% mixing rate, (m3/s);
 lam = 1/(60*60*24*365); % biological P consumption rate, 1/s
```

First set up the model parameters.

When you code the matrix manually, you enter row vectors in sequence. Each row is separated by semicolon (;)

% 3. set initial condition & time step dt = 60\*60\*24; % one day timestep N = 365\*10; % 10 year integration P = zeros(3,N); % preparing the solution storage Pa= zeros(3,N); % preparing the analytic solution storage Pb= zeros(3,N); % preparing the Euler backward storage P(:,1) = [2e-3; 2e-3; 2e-3]; % initial P conc. = uniform 2 mmolP/m3 Pa(:,1)= P(:,1); % initial cond for analytic sol. Pb(:,1)= P(:,1); % initial cond for euler backward sol. time=0:dt: (N-1)\*dt; % set up time array

In preparation for the numerical integration, prepare arrays for storing solutions. It makes the calculation faster.

```
% 4. execute numerical integration
G = eye(3,3) + dt*T; % Euler forward
H = inv(eye(3,3) - dt*T); % Euler backward
for n=1:(N-1)
    P(:,n+1) = G*P(:,n); % Euler step forward
    Pb(:,n+1) = H*Pb(:,n); % Euler backward
    Pa(:,n+1) = expm(T*time(n+1))*Pa(:,1); % analytic solution
end
```

This is the core of the model integration. We apply 3 different methods for comparison. Plotting script follows this section.

# The results



## Questions



This system tool approximately 3 years to reach the steady state.

Why 3 years?

Why did the value of deep ocean box hardly change?

## Why 3 years?

- Look at the eigenvalue
- Negative eigenvalues indicate the rate of exponential decay
- The inverse of eigenvalues gives the timescale that the initial condition decays.

>> [Q,D]=eig(T)		
Q =		
0.0098 0.0495 0.9987	0.9777 0.2088 -0.0235	-0.0027 -0.9999 0.0156
D =		
1.0e-07 >	k	
0	0	0
0	-0.3203	0
0	0	-0.3382

## Why 3 years?

- Look at the eigenvalue
- Negative eigenvalues indicate the rate of exponential decay
- The inverse of eigenvalues gives the timescale that the initial condition decays.



## Constant deep ocean?

- Look at the eigenvalue
- Negative eigenvalues indicate the rate of exponential decay
- The inverse of eigenvalues gives the timescale that the initial condition decays.

